

Characterization of the vertical structure of the atmosphere using ground-based remote sensing (III)

C. Yballa Hernández Pérez

PROJECT 3
(Call 2011)

Training in remote instrumental techniques (ground-based remote sensing) for detection and study of reactive gases and atmospheric aerosols

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Supervisors: Dr. Alberto Berjón Arroyo, Dr. Silvia Alonso Pérez

3rd October, 2014

Izana Atmospheric Research Center
CIAI

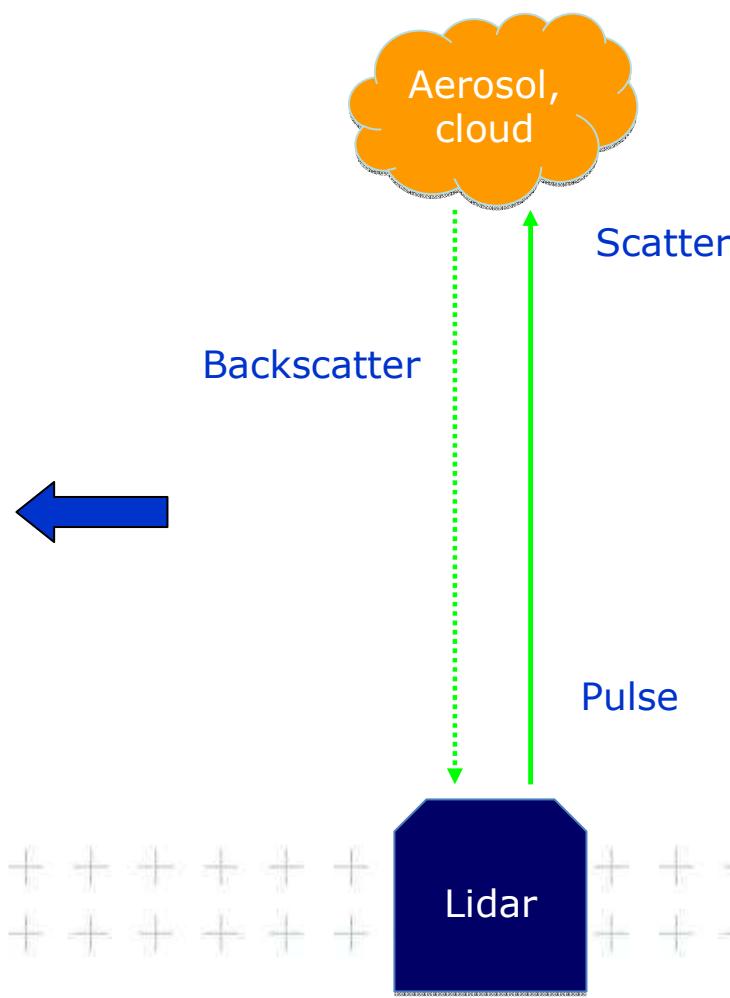
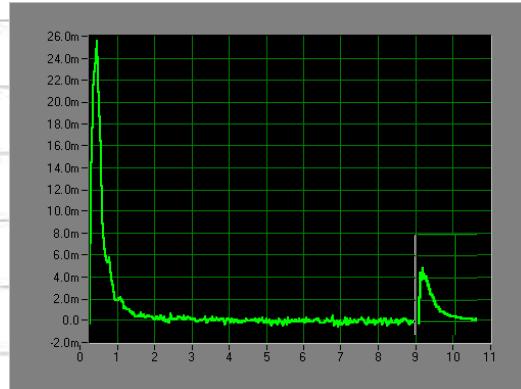


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 - Micro Pulse Lidar
 - Ceilometer
- Maintenance, calibration and evaluation of data
- Main results
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Introduction

- Lidar → capable of providing vertical profiles of aerosol and cloud structure



-Lidar



Micro Pulse Lidar (MPL-3)

- ✓ CIAI (AEMET) and INTA co-manage the MPL-3
- ✓ 523.5 nm
- ✓ 7 µJ
- ✓ 'eye-safe'



<http://mplnet.gsfc.nasa.gov>.



www.lidar.es/spalinet/es/



GALION

www.wmo.int/gaw/galion/index.html

-Ceilometer Vaisala CL51



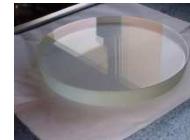
VAISALA

- ✓ Compact and robust lidar system
- ✓ Operates in extreme weather conditions
- ✓ 910 nm
- ✓ 3 µJ
- ✓ 'eye-safe'

Maintenance, calibrations and evaluation of data, MPL-3

Maintenance

- ✓ Temperature control : laser, telescope, detector, location...
- ✓ Humidity control
- ✓ Cleaning of optics
- ✓ Daily checking list



10%

Calibrations

- ✓ Darkcurrent (monthly)
- ✓ Afterpulse (monthly)
- ✓ Overlap (once a year)

20%

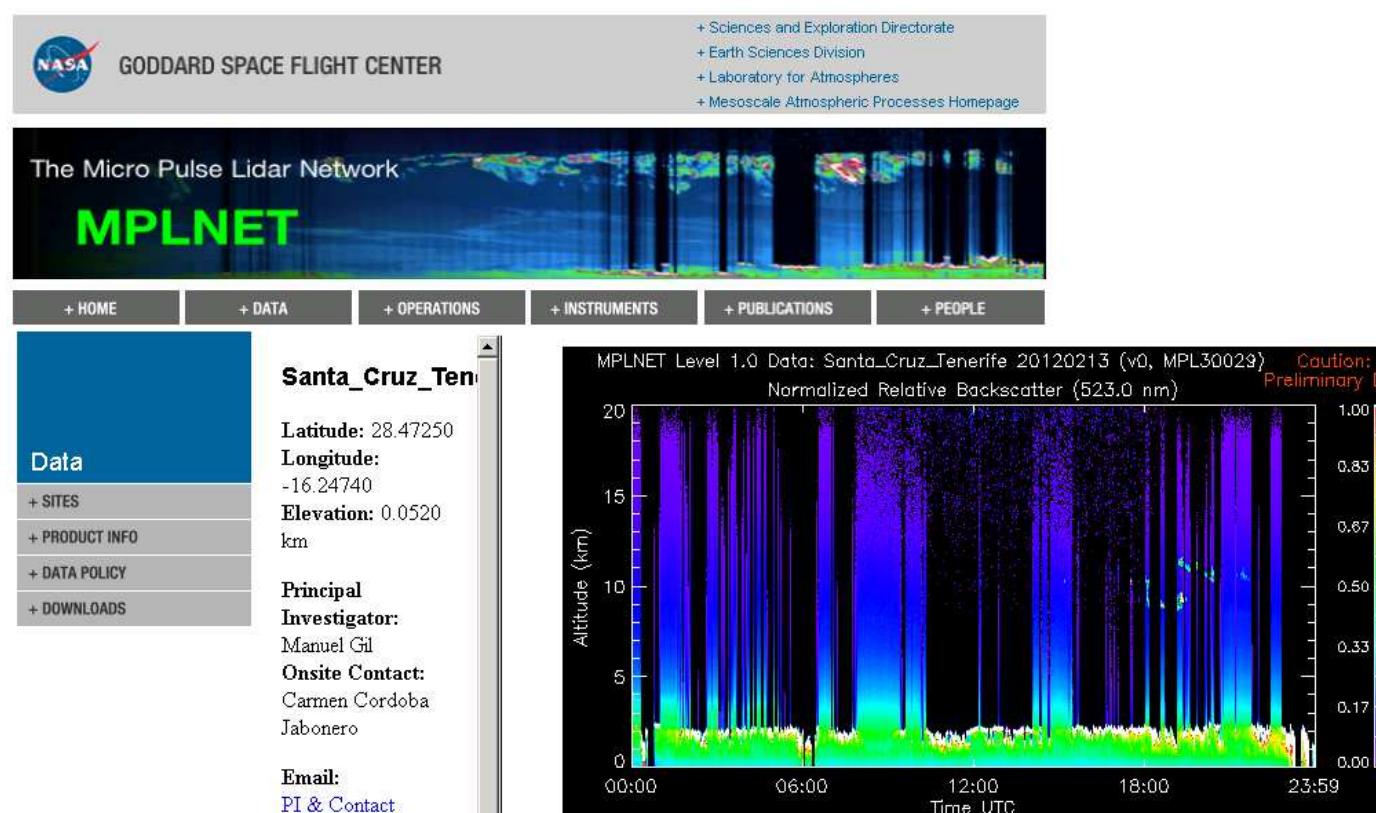
Data evaluation

- ✓ Calibrations
- ✓ Measurements

70%

Maintenance, calibrations and evaluation of data, MPL-3

Sending daily files of Lidar data to MPLNET



<http://mplnet.gsfc.nasa.gov/dat.htm>

Maintenance, calibrations and evaluation of data, Vaisala CL51

Maintenance

- ✓ Overall control of the instrument
- ✓ Cleaning of optics
- ✓ Daily checking list

40%

Calibrations

X

Data evaluation

- ✓ Measurements: height of clouds, BL, aerosols

60%

Main Results (1/3)

- New afterpulse (AFP) correction MPL-3

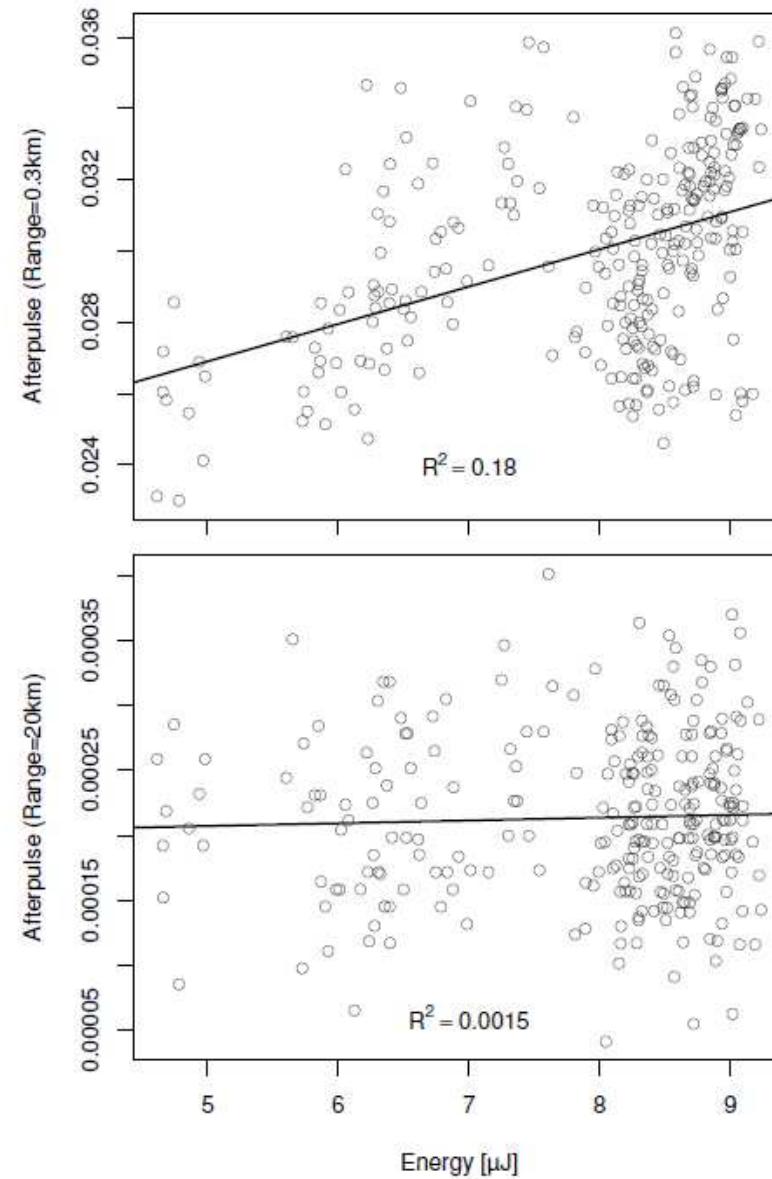
$$P(r) = r^2 \frac{n(r) - A(r) - D - B}{O(r)E}$$

$$A(r) = \frac{E}{E_0} A_0(r) \quad (\text{Welton and Campbell, 2002})$$

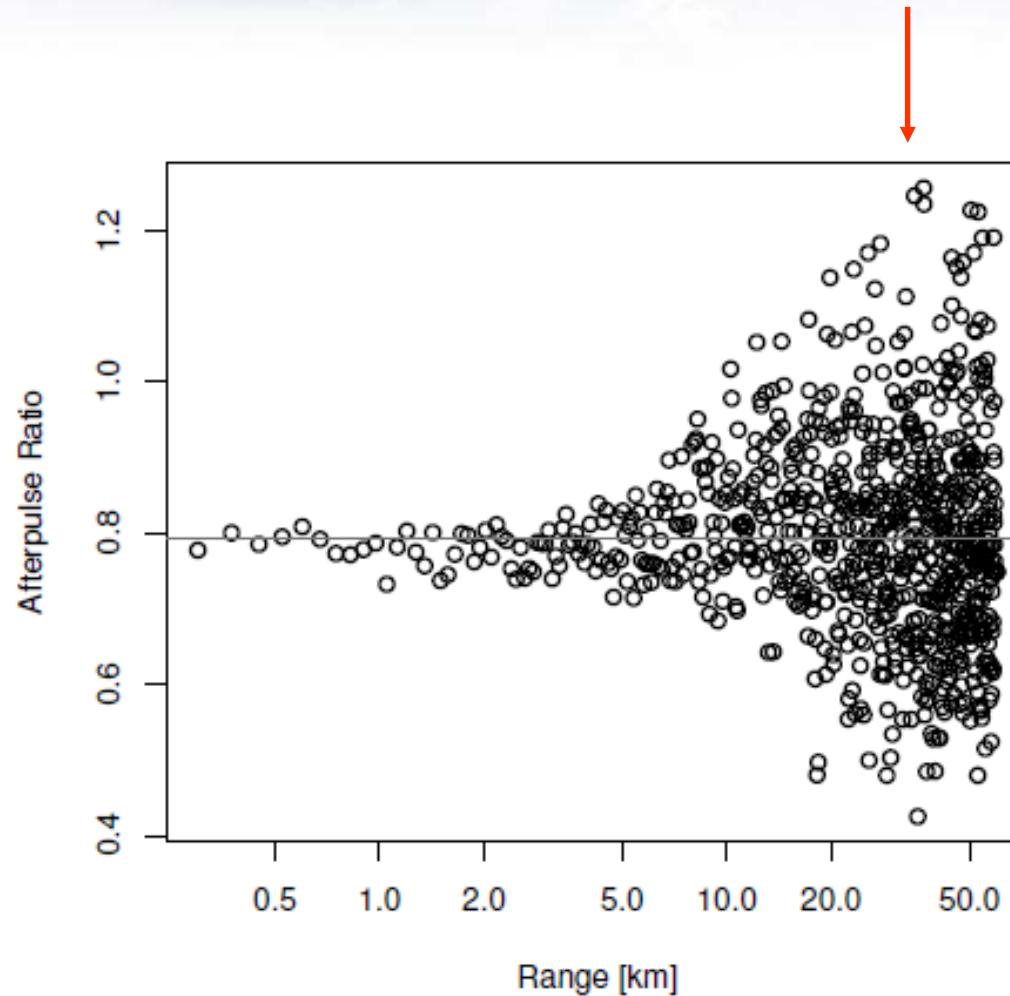
$$A(r) = F_r A_0(r) \quad (\text{CIAI})$$

- Dependence of the AFP on output pulse energy
- Dependence of the AFP on obstacle reflectance
- Effect of AFP correction on MPL background
- Effect of AFP correction on the analysis of aerosol intrusion at high altitudes

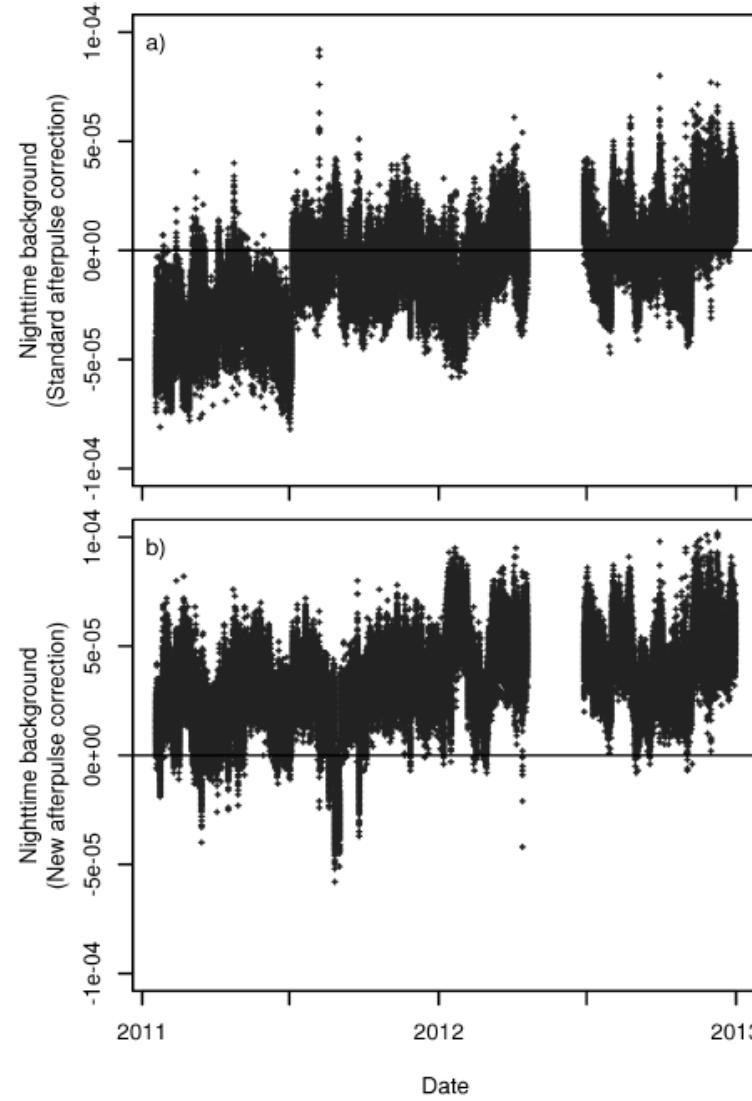
Dependence of the AFP on output pulse energy



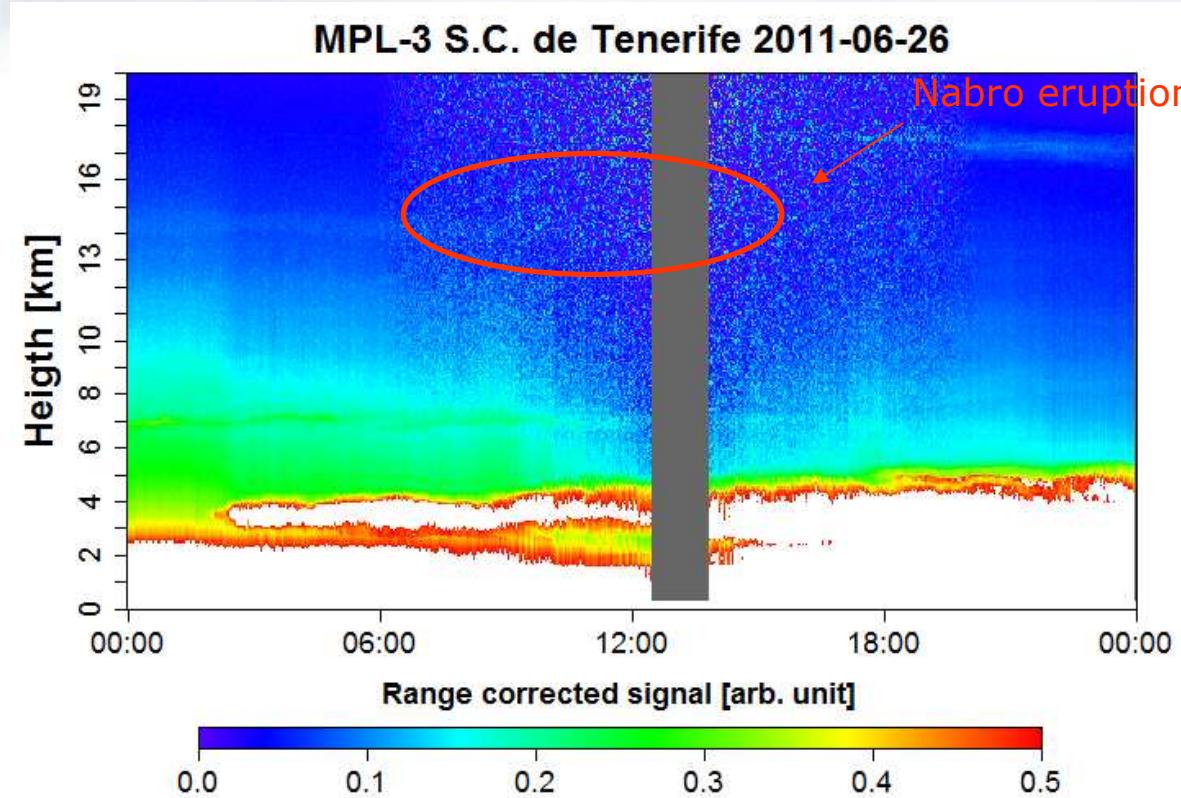
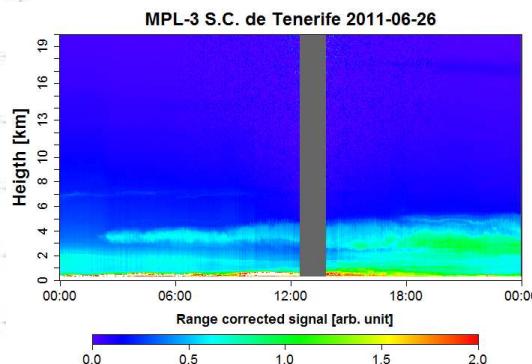
Dependence of the AFP on obstacle reflectance



Effect of AFP correction on MPL background



Effect of AFP correction on the analysis of aerosol intrusion at high altitudes



AFP_(Welton and Campbell; 2002)

AFP_(CIAI)

→ AOD=0.031+/-0.003 LR=80+/-6sr

→ AOD=0.021+/-0.002 LR=55+/-4sr

{
32% AOD
31% LR

Sawamura et al., 2012

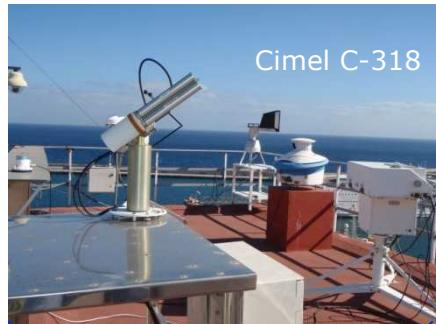
New afterpulse (AFP) correction MPL-3

Reprocessing of lidar data series → new algorithm

- Calibrations: darkcurrent, afterpulse, overlap
- Raw signal
- Range corrected signal
- Parameters: temperatures, energy,...
- Signal to noise ratio
- Mapping

Main Results (2/3)

- Inversion algorithm (Fernald 1984; Klett ,1985) “one layer”



$$P_{NRB}(r) = C(\beta_M(r) + \beta_P(r)) e^{-2 \int_0^z \sigma_M(r') dr'} e^{-2 \int_0^z \sigma_P(r') dr'}$$

- 3 unknowns
- To solve the problem:

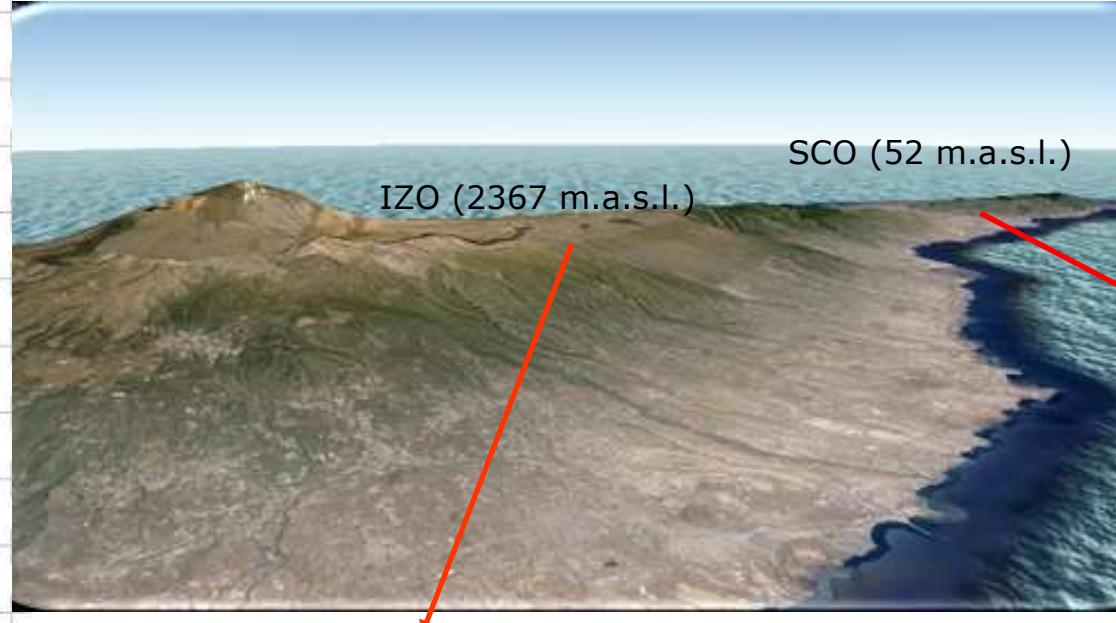
$Z_{ref} \rightarrow \times$

$LR = \sigma/\beta$

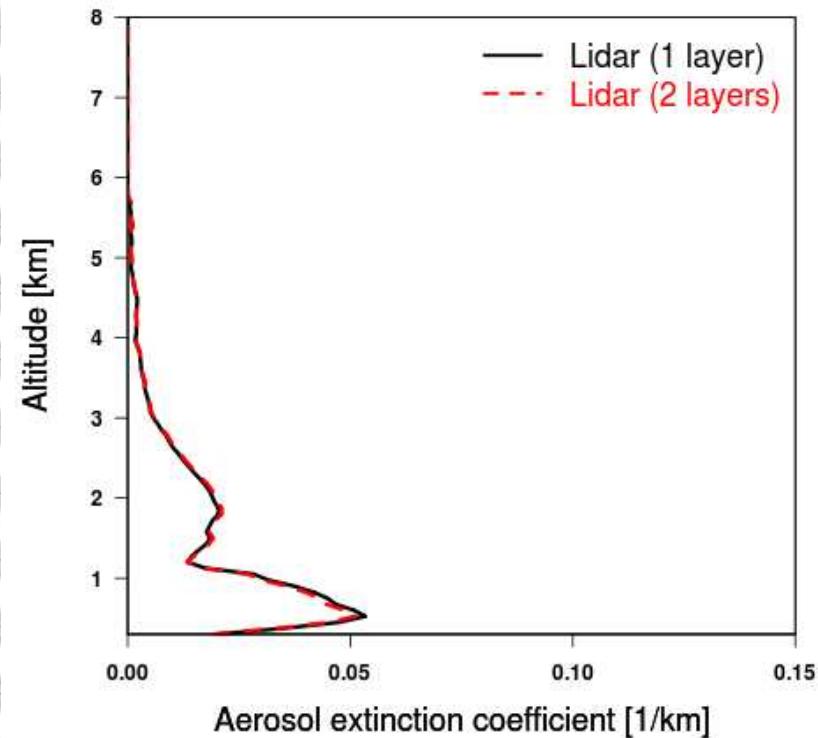
$AOD_{MPL} \approx AOD_{AERONET}$

Main Results (3/3)

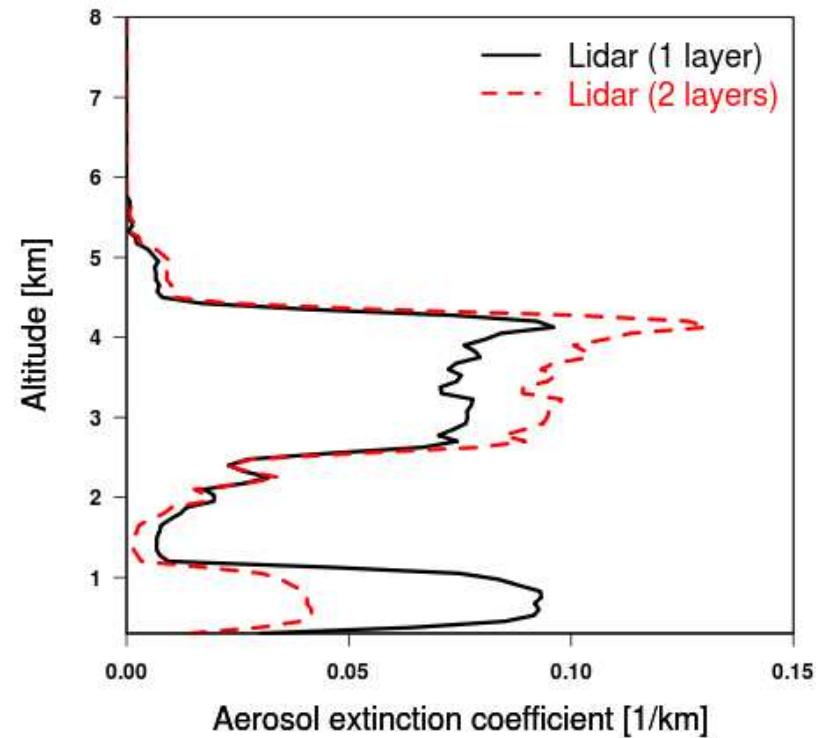
- Inversion algorithm (Fernald 1984; Klett ,1985; Sasano,1985) “two layers”



MPL-3 SCO (2012-03-14)



MPL-3 SCO (2012-07-09)



Inversion	LR(sr)
One layer	24
Two layers	23 25

Marine	(LR=20-30 sr)
Mixed	(LR=30-45 sr)
Dust	(LR= 45-70 sr)

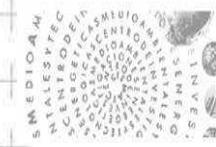
Inversion	LR(sr)
One layer	45
Two layers	21 63

Müller et al., 2007 → FT LR=29 sr

Publications and Congress Communications, 2011-2014 (1/5)



AEMET
Agencia Estatal de Meteorología



RECTA 2011

V Reunión Española de Ciencia y Tecnología de Aerosoles
Madrid, 27 - 29 de Junio de 2011, CIEMAT



4-9 Sept 2011

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Synergistic monitoring of Saharan dust plumes and potential impact on surface: a case study of dust transport from Canary Islands to Iberian Peninsula

C. Córdoba-Jabonero, J. Sorribas, J. L. Guerrero-Rascado, J. A. Adame, Y. Hernández, H. Lyamani, V. Cachorro, M. Gil, L. Alados-Arboledas, E. Cuevas, and B. de la Morena, Instituto de Estudios Atmosféricos (IEA), Consejo Superior de Investigaciones Científicas (CSIC), Madrid, Spain

Abstract: The complex set of meteorological advection processes involved in the transport of Saharan dust plumes to the Iberian Peninsula makes their detection and identification difficult. In this work we have used two different methods to detect and identify Saharan dust plumes during the African dust event of June 2010. One method is based on the vertical structure analysis of the dust plume using lidar and ceilometer data, and the other is based on the synergistic use of the information provided by the optical remote sensing instruments (ceilometer, lidar, and micro-pulse lidar) and the chemical tracers (PM2.5, mineral dust, and organic matter) measured at the IEA station located in Madrid. The lidar and ceilometer data were used to detect the presence of Saharan dust plumes and to obtain information about their vertical structure. The chemical tracers were used to identify the plumes and to determine their origin. The results obtained show that the lidar and ceilometer data are able to detect the presence of Saharan dust plumes during the African dust event of June 2010, and that the synergistic use of the lidar and ceilometer data and the chemical tracers is able to identify the plumes and to determine their origin. The results obtained show that the lidar and ceilometer data are able to detect the presence of Saharan dust plumes during the African dust event of June 2010, and that the synergistic use of the lidar and ceilometer data and the chemical tracers is able to identify the plumes and to determine their origin.

Keywords: Synergistic monitoring, Saharan dust plumes, lidar, ceilometer, chemical tracers, African dust event, Iberian Peninsula, Canary Islands.

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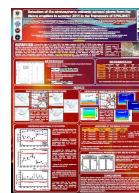
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Publications and Congress Communications, 2011-2014 (2/5)



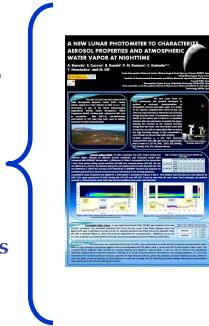
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Atmos. Meas. Tech., 5, 5527–5569, 2012
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A new method for nocturnal aerosol measurements
with a lunar photometer prototype

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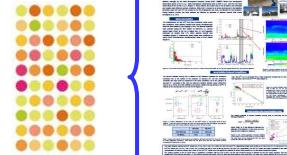
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This paper present the preliminary results of nocturnal Aerosol Optical Depth (τ_a) and Angstrom Exponent (α) measurements made with a prototype of the instrument named Cimel CE-318T. Due to the variation of the moon's illumination inherent to the lunar cycle, the typical Langley Method cannot be applied. In this paper, we propose three different methods to estimate the aerosol optical properties. In order to validate the results, we have selected three events which encompass several nights and sea days under different atmospheric conditions and different dust intrusion episodes. Method1 is introduced in this work as a modification of the usual Langley Method. This semi-empirical irradiance from a lunar irradiance model, provides reliable results for the aerosol optical properties. Method2 consists of transforming the current calibration coefficients of the Cimel CE-318T into a new set of coefficients that will represent one of the largest uncertainties in climate change studies (IPCC, 2007). The high uncertainty associated with the role played by aerosols in radiative forcing is due to the lack of validation of satellite-based aerosol climatology. In this sense, the Aerosol Robotic Network (AERONET) is considered one of the most promising datasets for validation of satellite-based aerosol retrievals (Kaufman et al., 1997). The aerosol optical depth (τ_a) at a certain wavelength is the standard parameter used to describe the extinction of light in the atmosphere. AERONET Spectral dependence of τ_a is mainly driven by the scattering efficiency and can be expressed by means of the Angstrom exponent (α). The Angstrom exponent (α) in the solar spectrum, the Angstrom exponent (α) is a good indicator of the dominant size of the atmospheric particles. τ_a and α are related by the equation $\alpha = (\log \tau_a)/(\log \lambda)$, which can be embedded in the error propagation in Method1. The good results obtained from the comparison against a second CE-318T instrument, provide independent and transitive validation to satellite-based

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Barreto, A., E. Cuevas, B. Damiri, C. Guirado, T. Berkoff, A. J. Berjón, Y. Hernández, F. Almansa, and M. Gil, A new method for nocturnal aerosol measurements with a lunar photometer prototype, Atmos. Meas. Tech. Discuss., 5, 5527-5569, doi:10.5194/amt-5-5527-2012, 2012.

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Publications and Congress Communications, 2011-2014 (5/5)



The MACC-II 2007-2008 Reanalysis: Atmospheric Dust Evaluation and Characterization over Northern Africa and Middle East

E. Cuevas¹, C. Camino^{1,2}, A. Benedetti³, S. Basart⁴, E. Terradellas⁵, J.M. Baldasano^{2,4}, J.-J Morcrette⁶, B. Marticorena⁶, P. Goloub⁷, A. Mortier⁷, A. Berjón¹, Y. Hernández¹, M. Gil-Ojeda⁸, M. Schulz⁹

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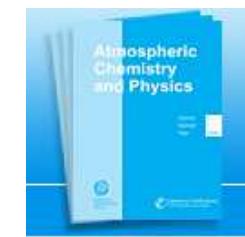
⁷Laboratoire d'Optique Atmosphérique, Université Lille 1, Lille, France

⁸Atmospheric Research and Instrumentation Branch, INTA, Madrid, Spain

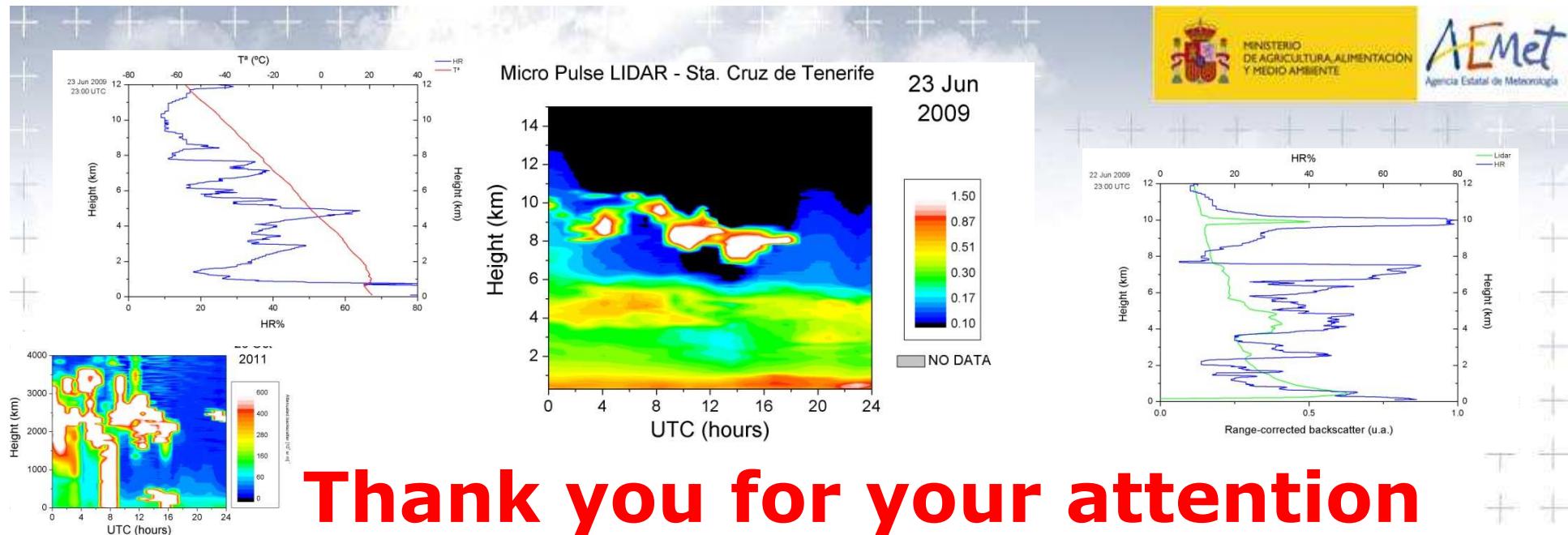
⁹Climate and Air pollution Section, Norwegian Meteorological Institute, Oslo, Norway

Abstract

In the present work, atmospheric mineral dust from a MACC-II short reanalysis run for two years (2007-2008), has been evaluated over Northern Africa and Middle East using satellite aerosol products (from MISR, MODIS and OMI satellite sensors), ground-based AERONET data, in-situ PM10 concentrations from AMMA, and extinction vertical profiles from two ground-based lidars and CALIOP. The MACC-II aerosol optical depth (AOD) spatial and temporal (seasonal and interannual) variability shows good agreement with those provided by satellite sensors. The capability of the model to reproduce AOD, Ångström exponent (AE) and dust optical depth (DOD) from daily to seasonal time-scale is quantified over twenty-six AERONET stations located in eight geographically distinct regions by using statistical parameters. Overall DOD seasonal variation is fairly well simulated by MACC-II in all regions, although the correlation is significantly higher in dust transport regions than in dust source regions. The ability of MACC-II in reproducing dust vertical profiles has been assessed by comparing seasonal averaged extinction vertical profiles simulated by MACC-II under dust conditions with



Atmospheric Chemistry and Physics



Thank you for your attention

